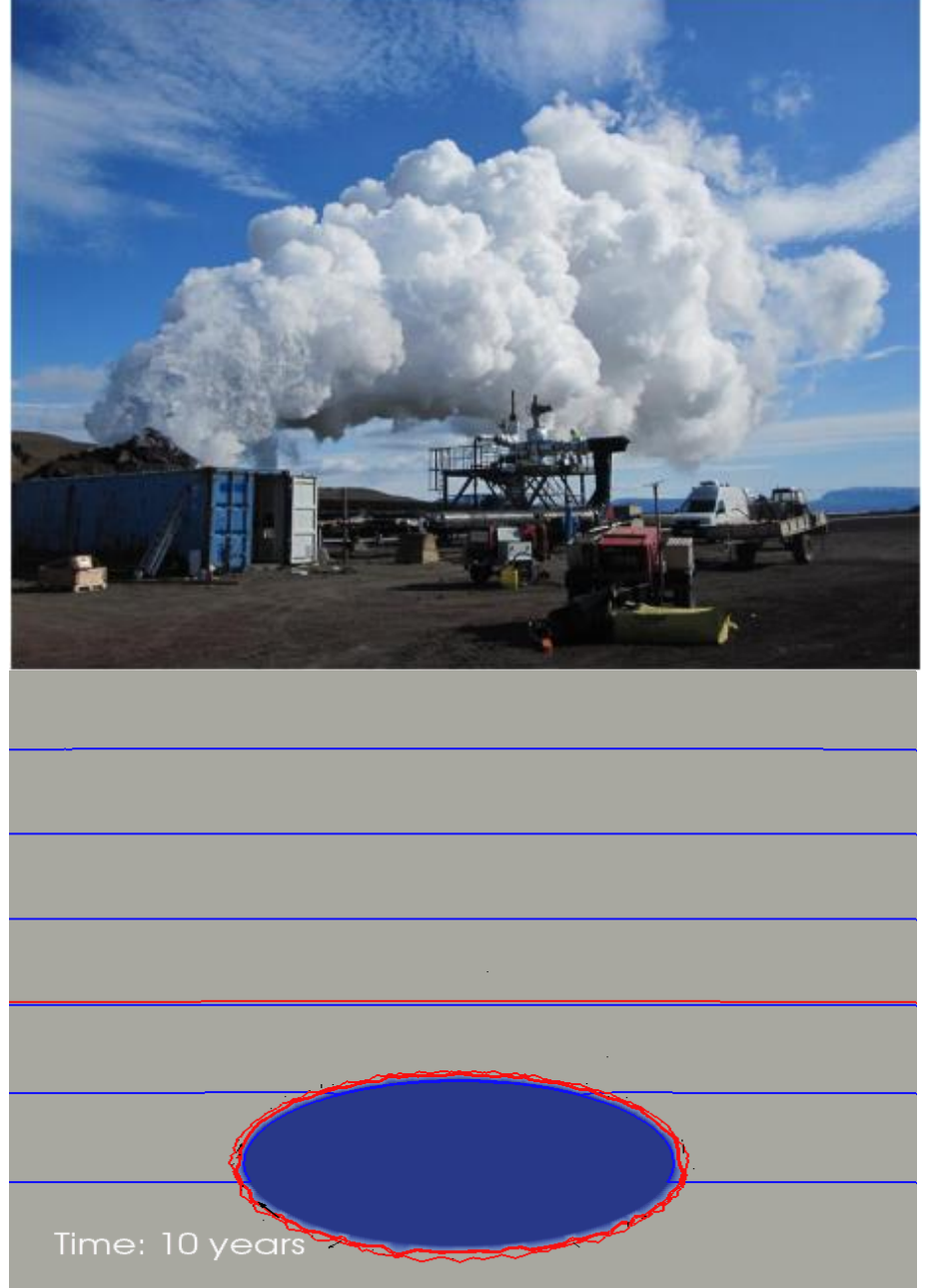


Modeling the evolution of volcanic geothermal systems with CSMP++

Samuel Scott
Reykjavik University

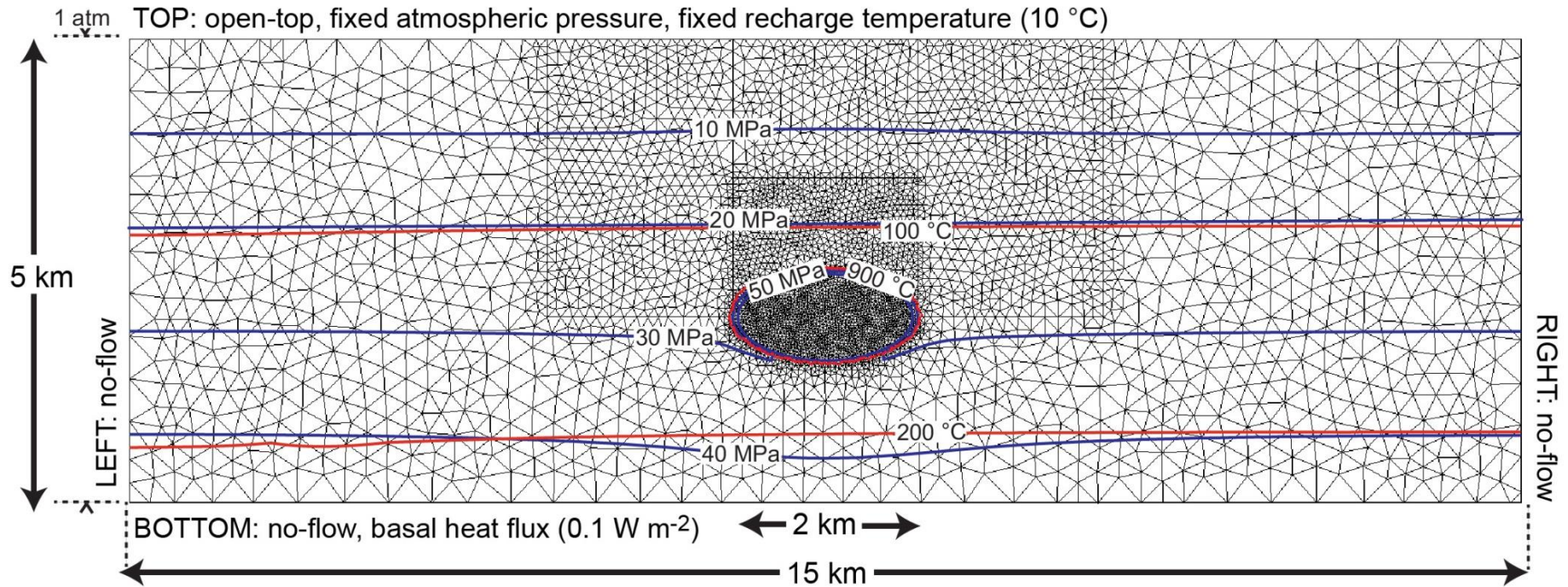
Special thanks to:
Thomas Driesner (ETH Zurich)
Philipp Weis (GFZ Potsdam)



Outline:

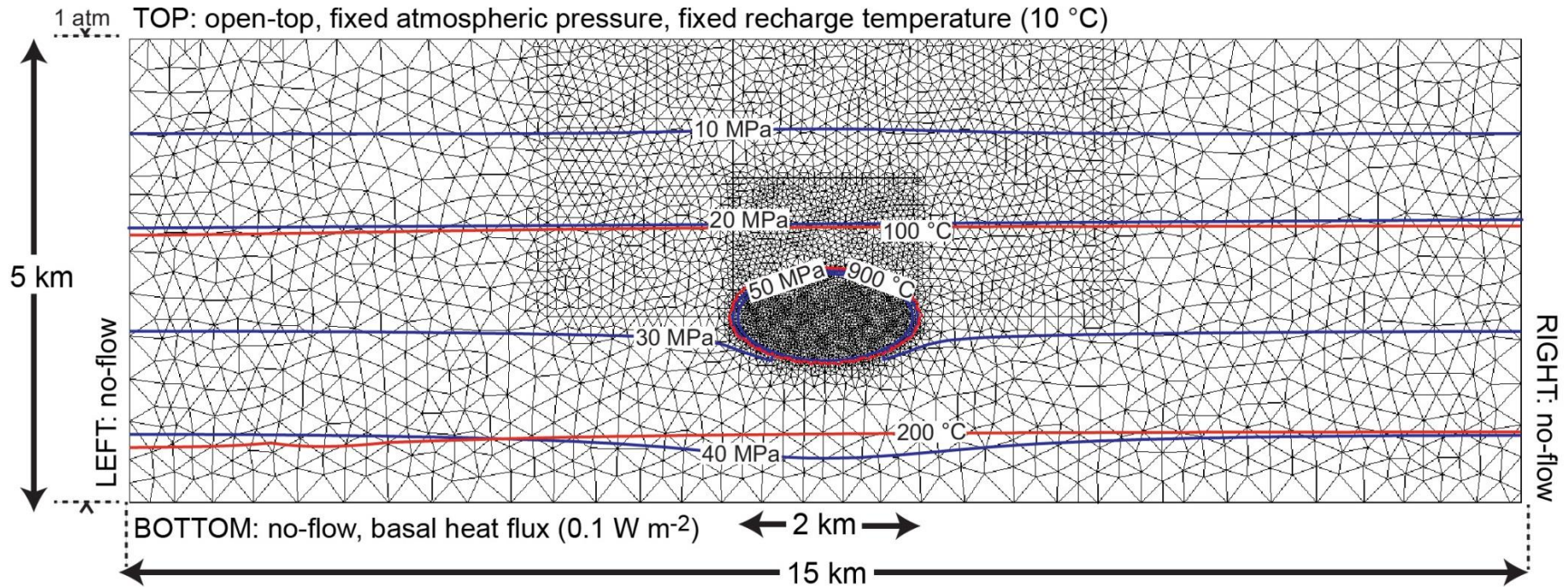
- Include a magmatic heat source in hydrologic models
- Simulate formation of supercritical geothermal resources
 - Published in *Nature Communications* (2015)
- Investigate how geology controls thermal structure
 - Published in *Geothermics* (2016)
- Compare meteoric- and seawater-recharged systems
 - Published in *Geophysical Research Letters* (2017)

Model set-up



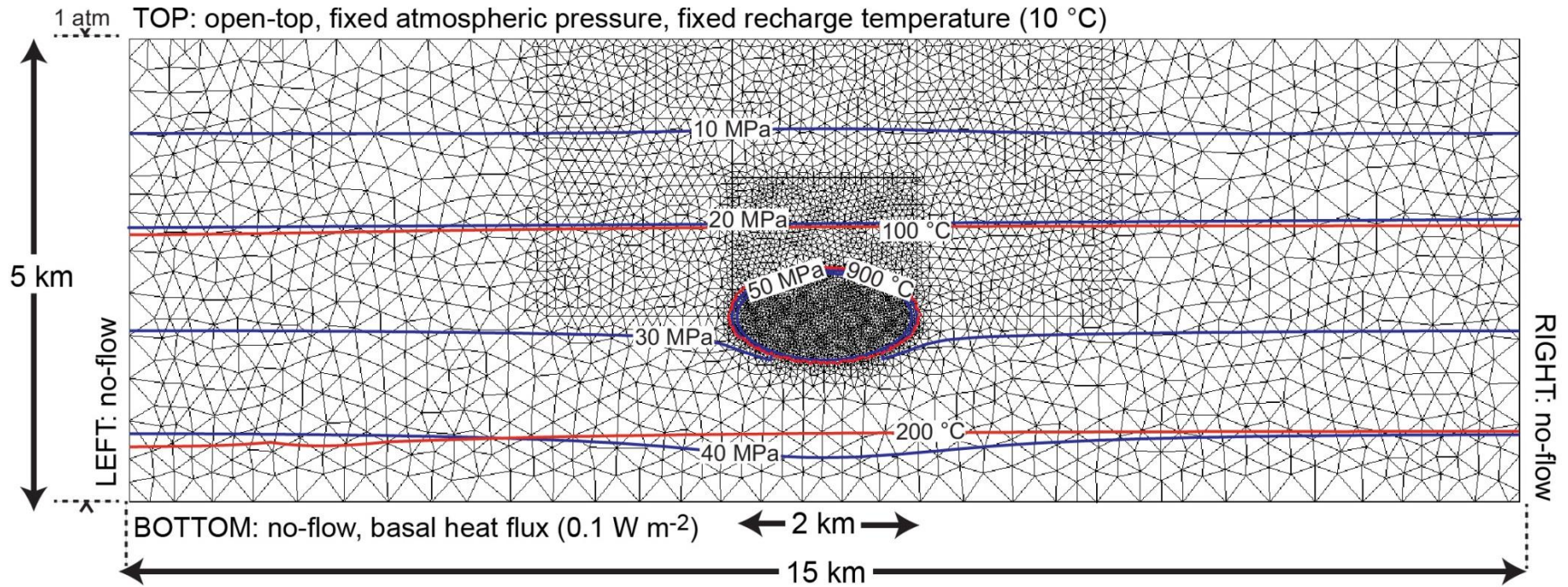
- Complex Systems Modeling Platform (CSMP++)
 - 2D porous media
 - Homogenous/isotropic permeability
 - Temperature-dependent permeability (T_{BDT})

Model set-up



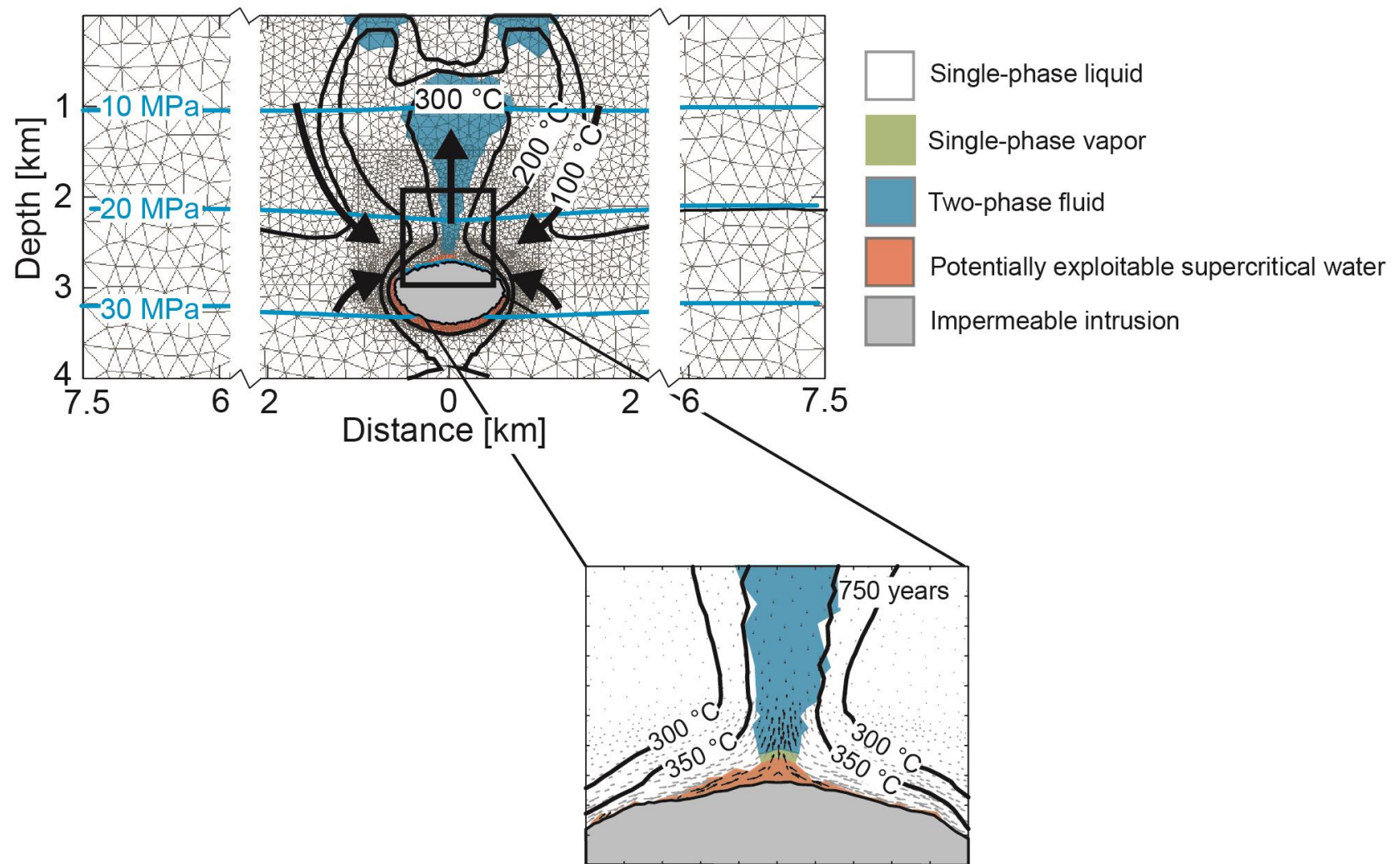
- Governing equations
 - Two-phase Darcy's law
 - Mass conservation
 - Energy conservation (enthalpy-based)

Model set-up



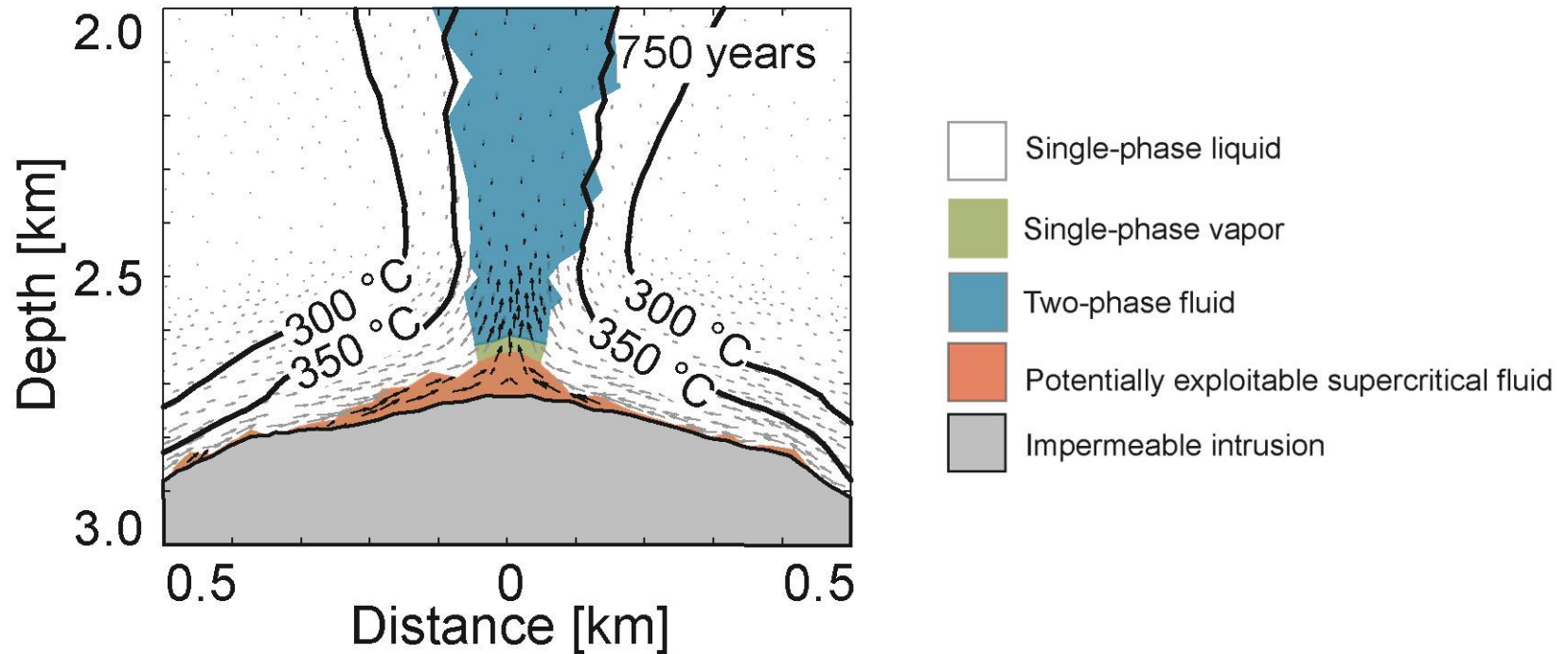
- Control Volume-Finite Element Numerical Scheme

Example results - high permeability



Scott et al., 2015, *Nature Communications*

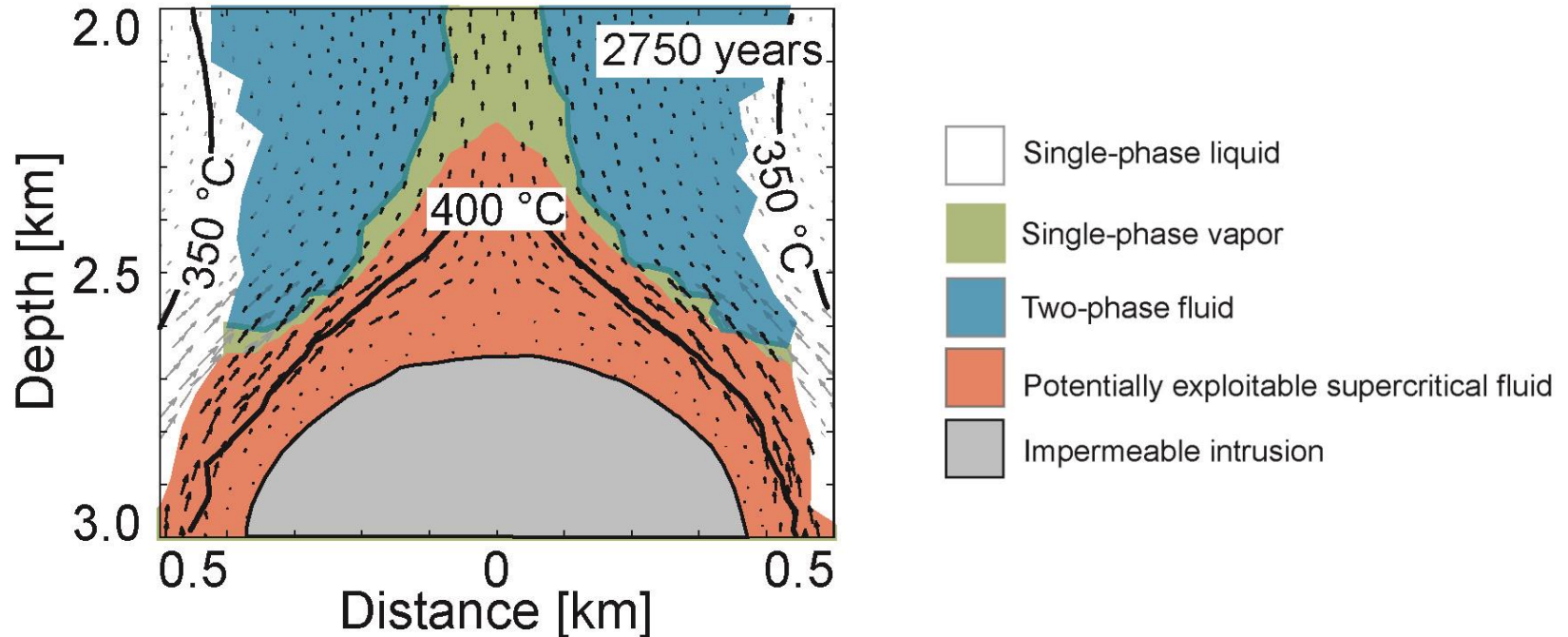
High permeability, $T_{BDT} = 450 \text{ }^\circ\text{C}$



- ‘Potentially exploitable supercritical resources’ defined as:
 - $T > 373.9 \text{ }^\circ\text{C}$, $h > 2.086 \text{ MJ/kg}$, $k > 10^{-16} \text{ m}^2$

Scott et al., 2015, *Nature Communications*

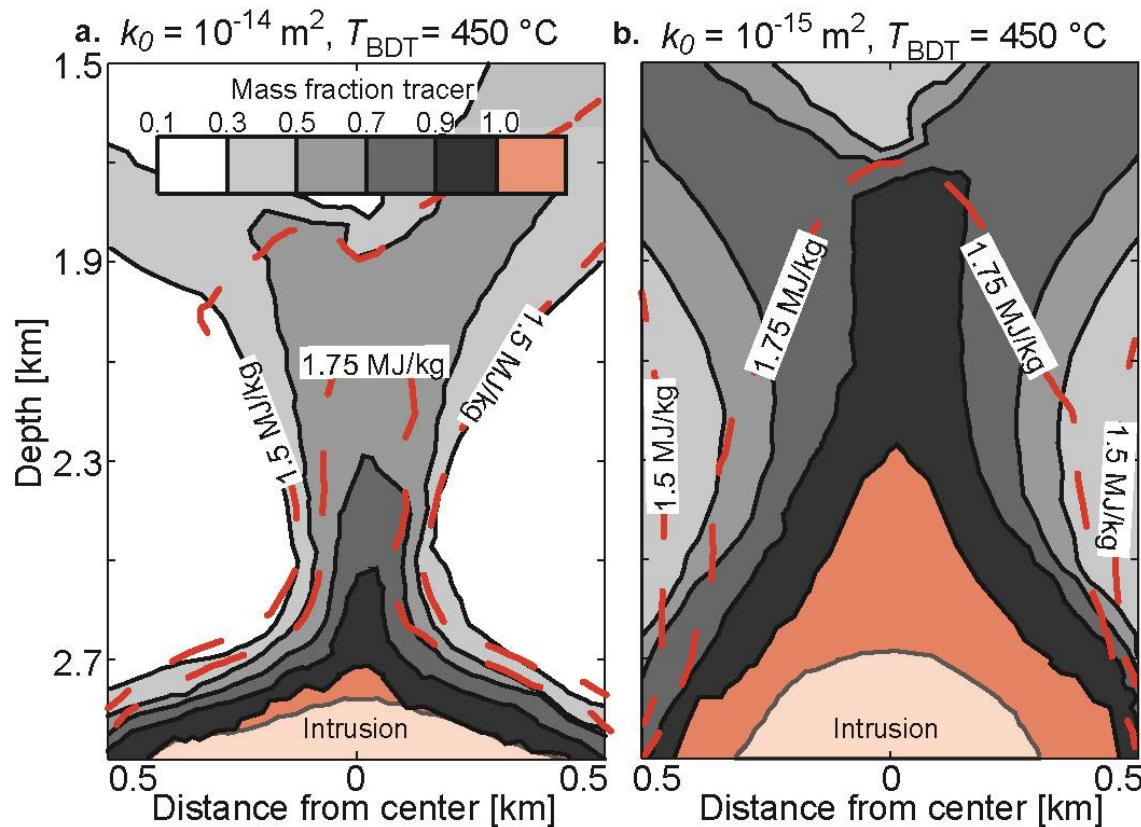
Intermediate permeability, $T_{BDT} = 450 \text{ }^\circ\text{C}$



- Supercritical resources much larger in spatial extent and show hotter temperatures

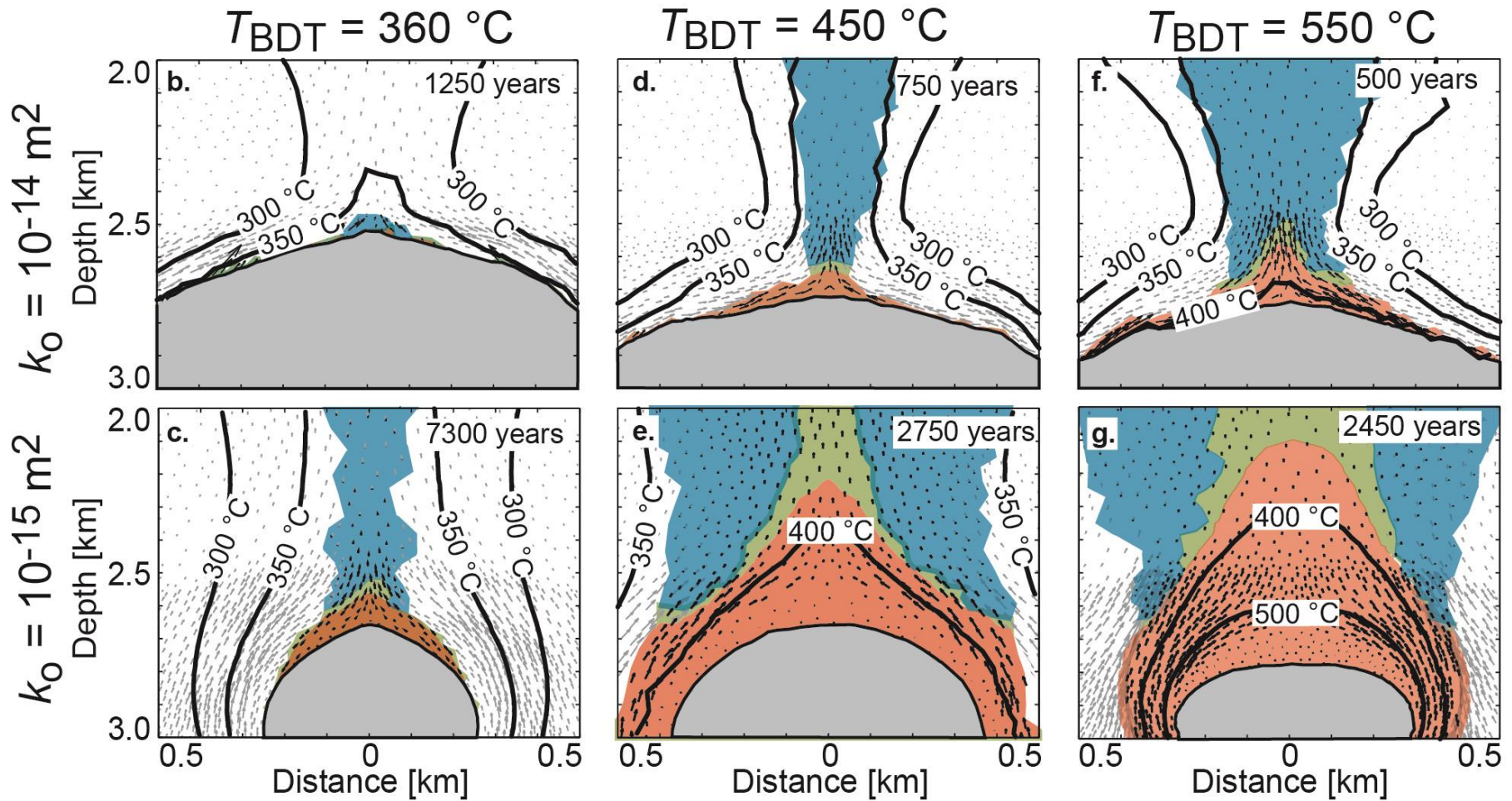
Scott et al., 2015, *Nature Communications*

Permeability controls fluid mixing



Conventional geothermal resources result from mixing of ascending supercritical and cooler circulating waters

Scott et al., 2015, *Nature Communications*

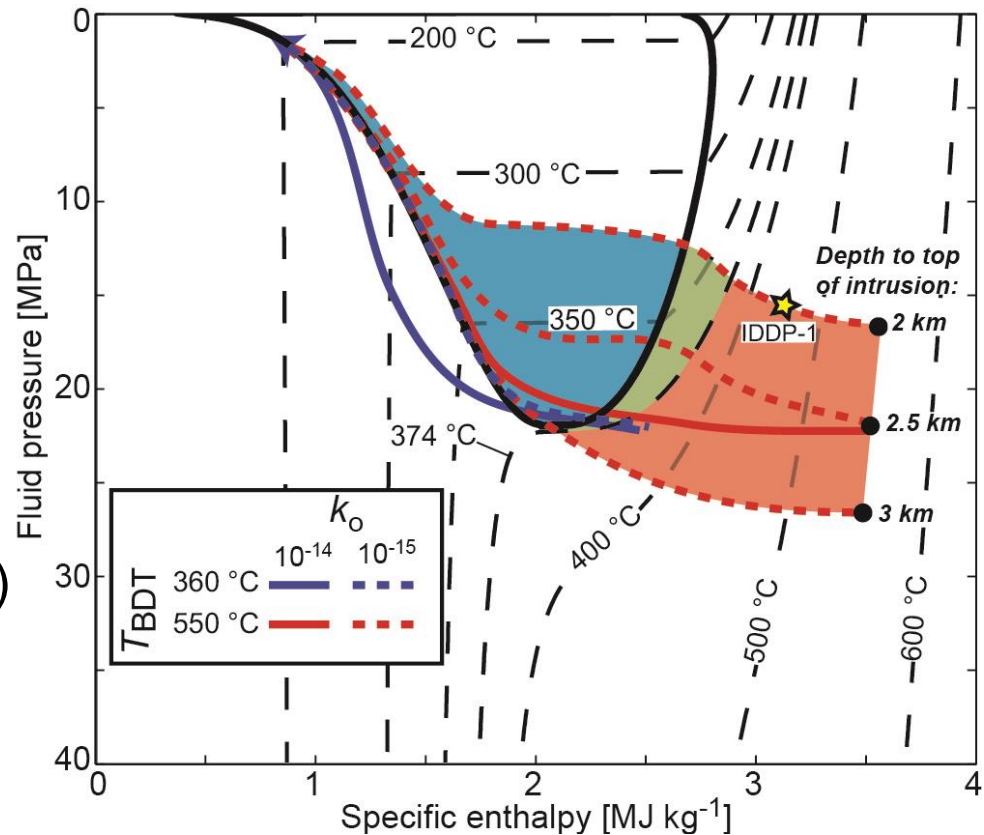


Scott et al., 2015, *Nature Communications*

Conclusions: Meteoric water systems

Larger, hotter supercritical zones ($T > 374\text{ °C}$, $h > 2.083\text{ MJ kg}^{-1}$) favoured by:

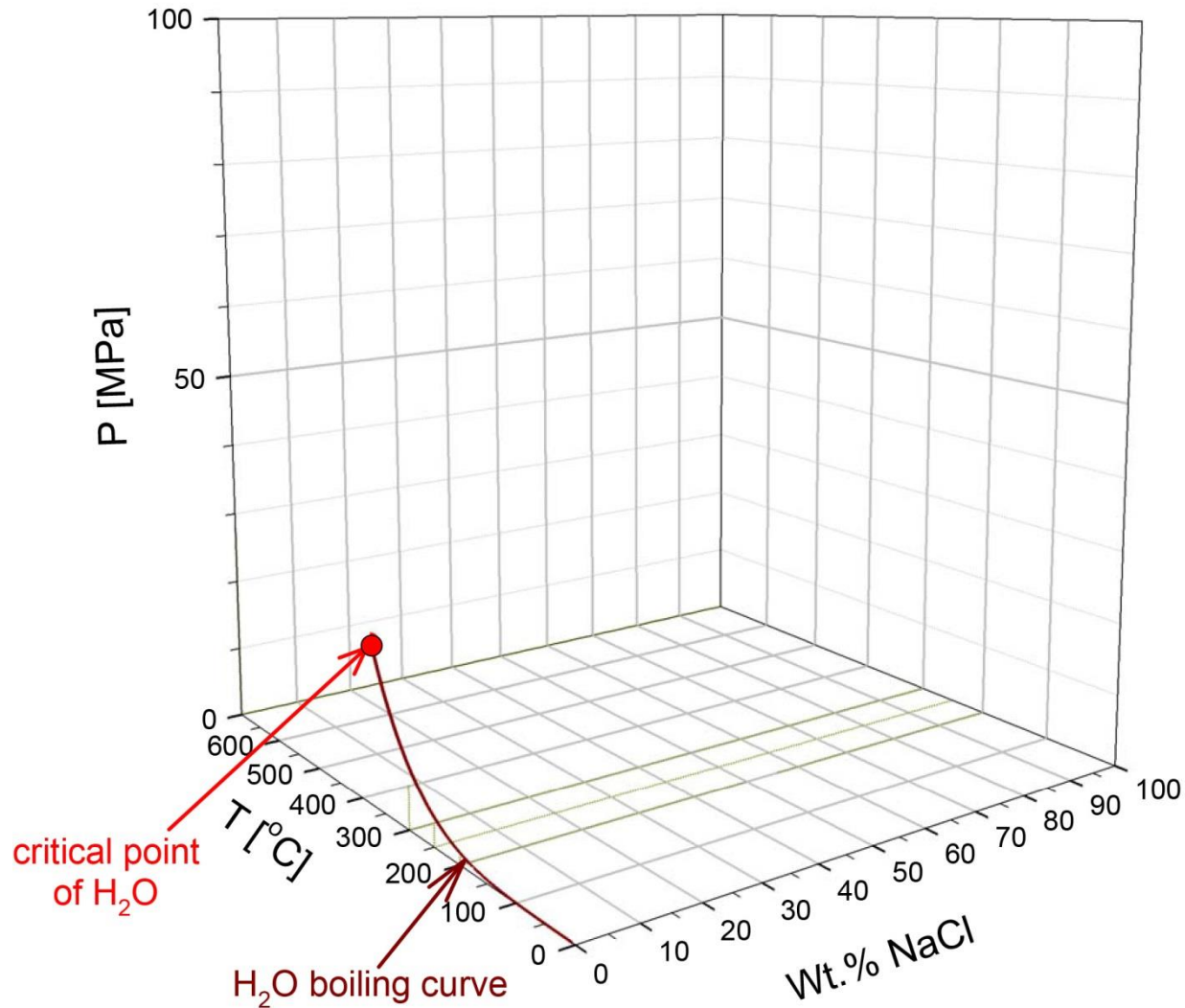
- Permeability near 10^{-15} m^2
- High T_{BDT} ($>400\text{ °C}$)
- Shallow intrusion depth (2-3 km)



- **IDDIP-1:** Measured reservoir conditions match predicted values assuming appropriate values for the geologic controls

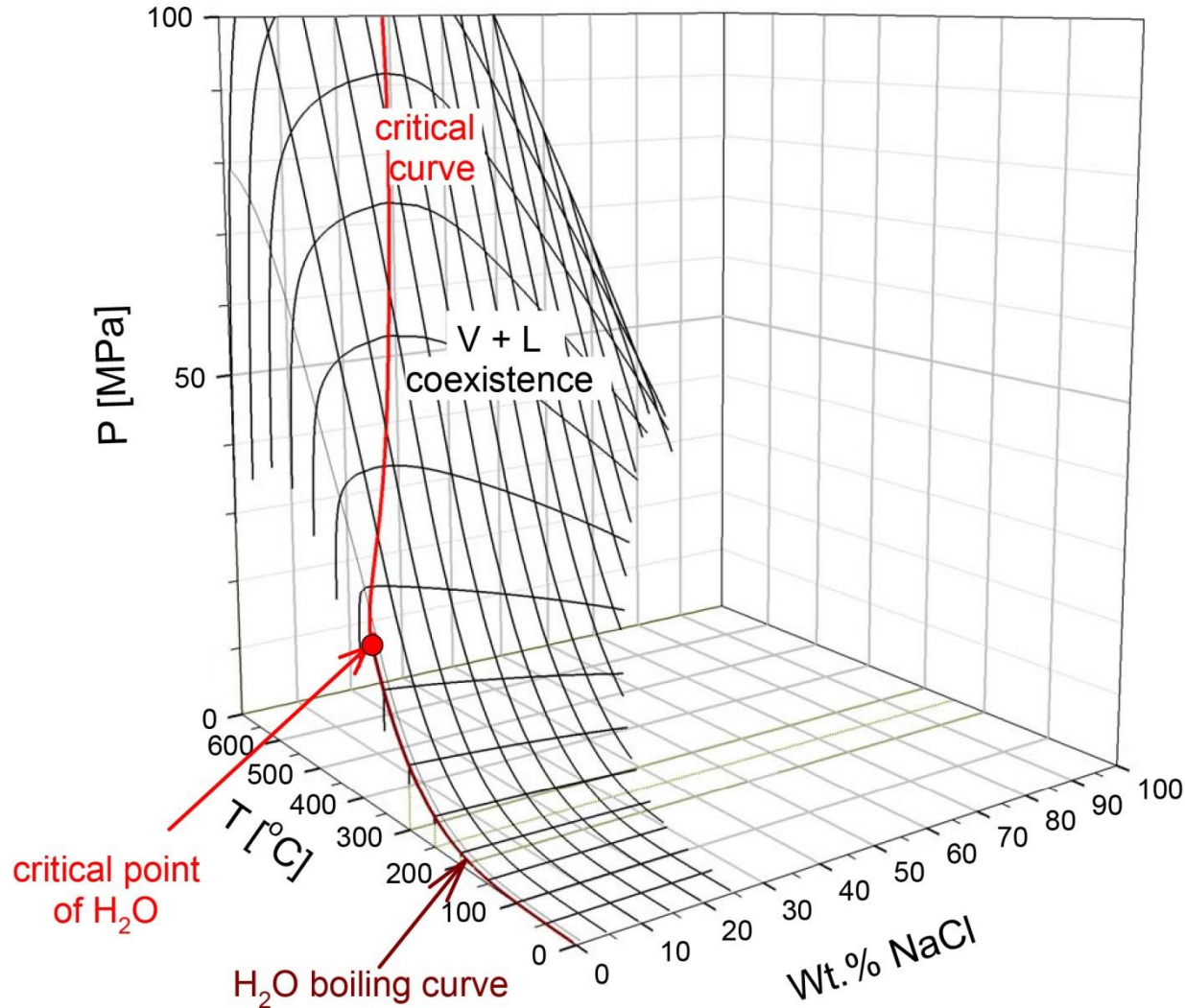
Scott et al., 2015, *Nature Communications*

The system $\text{H}_2\text{O}-\text{NaCl}$



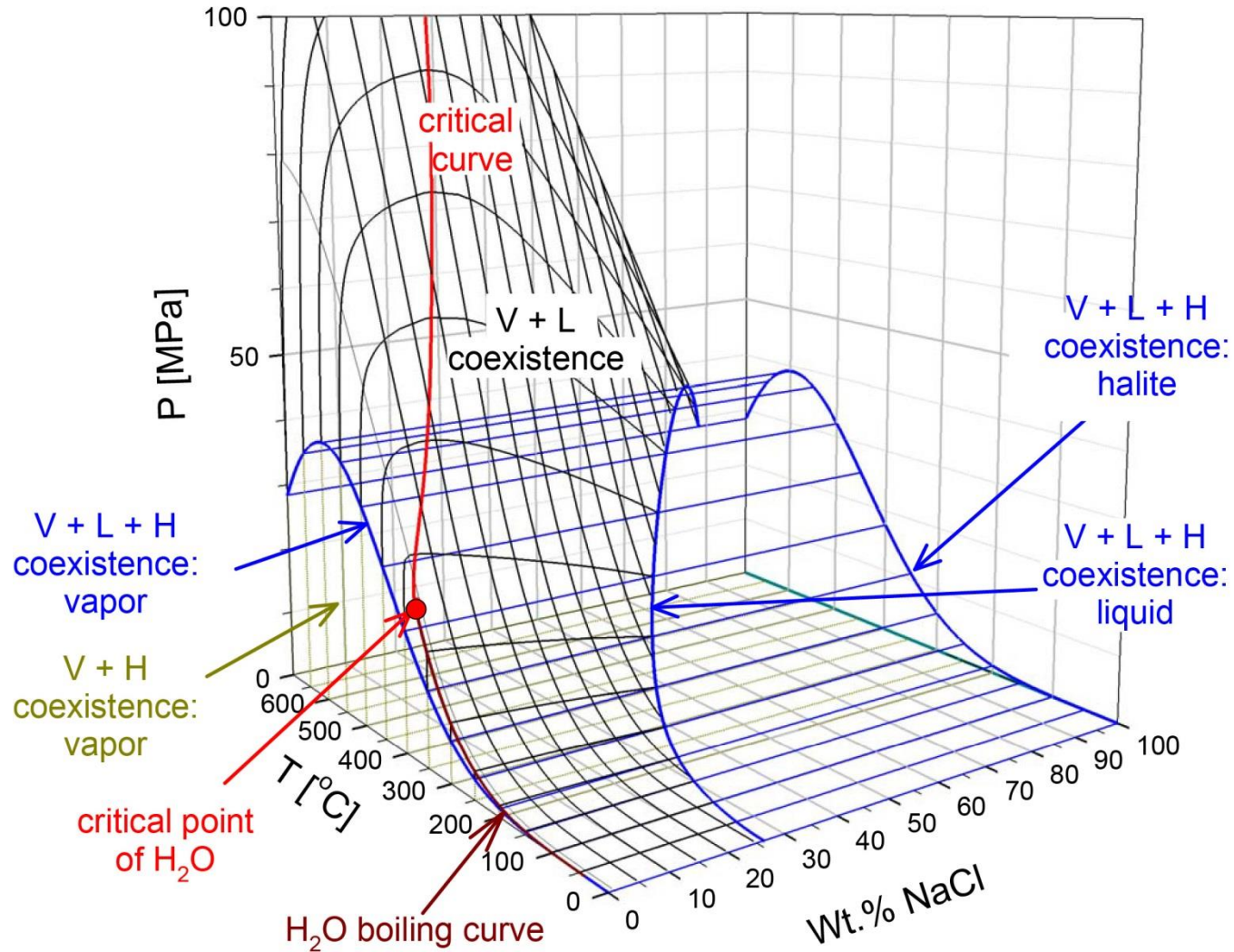
Driesner & Heinrich, GCA, 2007

The system $\text{H}_2\text{O}-\text{NaCl}$



Driesner & Heinrich, GCA, 2007

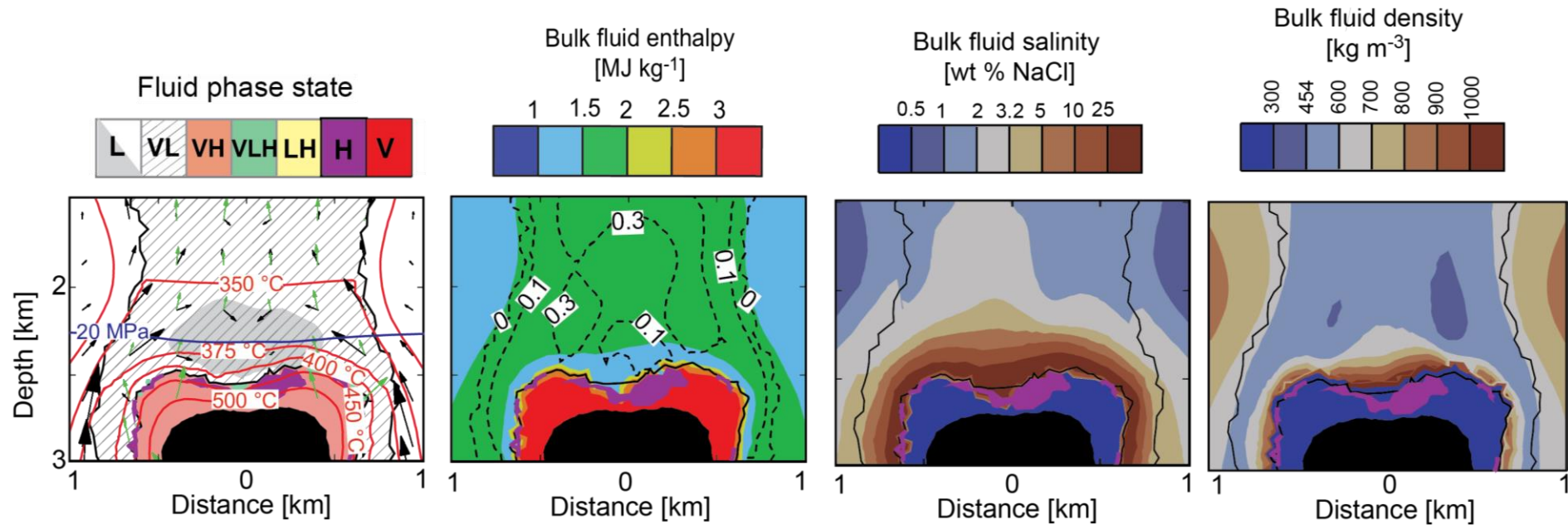
The system H₂O-NaCl



Driesner & Heinrich, GCA, 2007

Shallow intrusion -

Initial salinity 3.2 wt %, $k_0 = 10^{-15} \text{ m}^2$, $T_{BDT} = 550 \text{ }^\circ\text{C}$

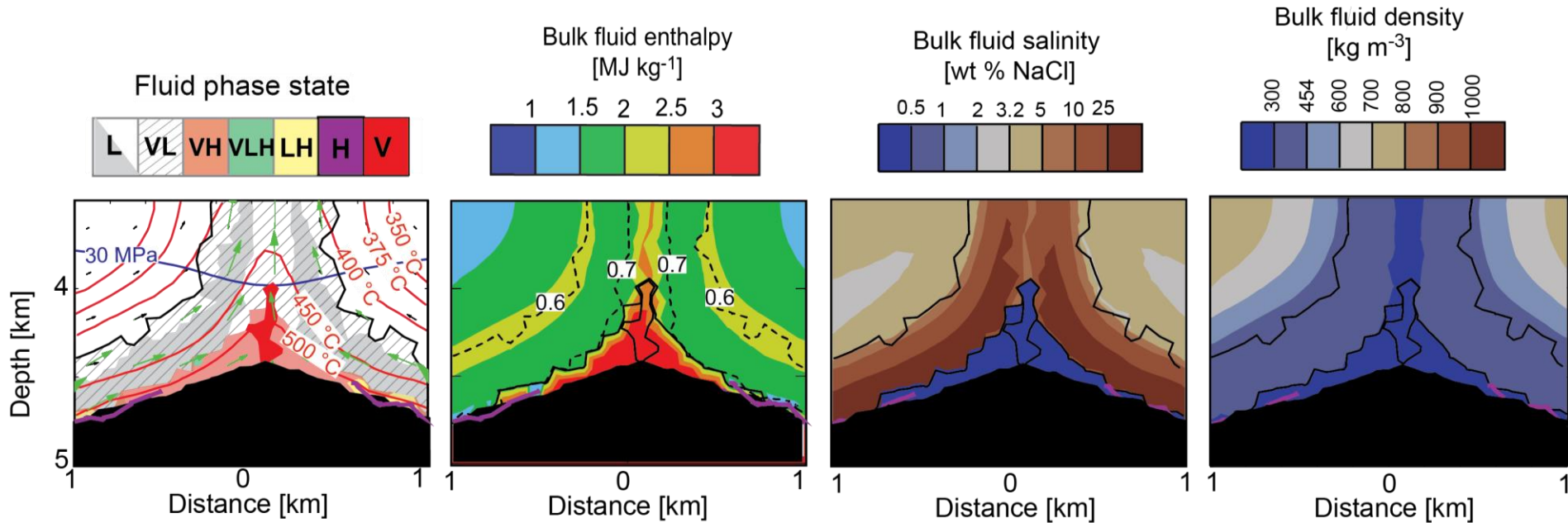


Systems develop deep hypersaline brines
underlain by halite-filled pore space

Scott et al., 2017, *Geophysical Research Letters*

Deep intrusion -

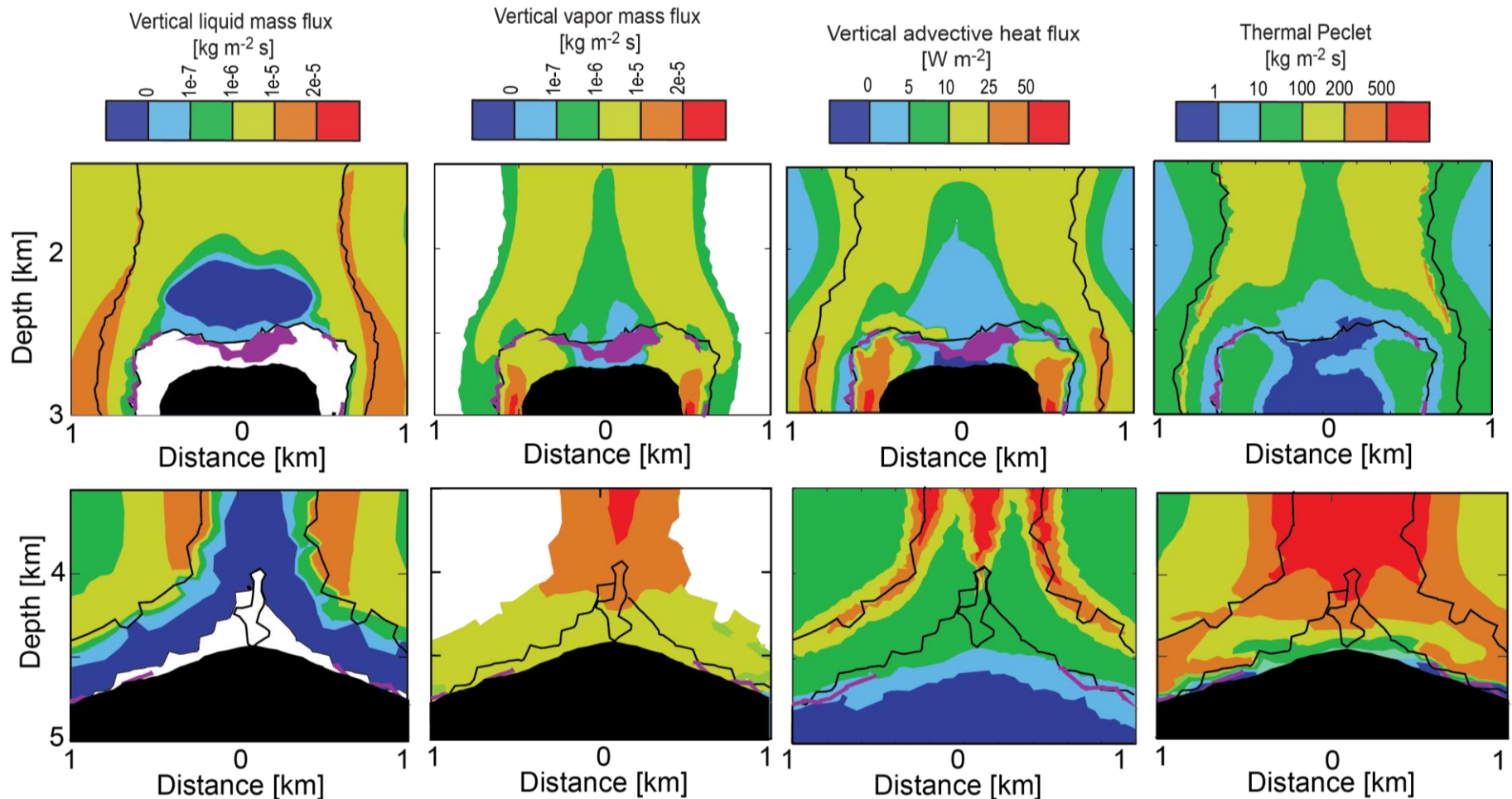
Initial salinity 3.2 wt %, $k_0 = 10^{-15} \text{ m}^2$, $T_{BDT} = 550 \text{ }^\circ\text{C}$



Halite-undersaturated vapor above center of intrusion

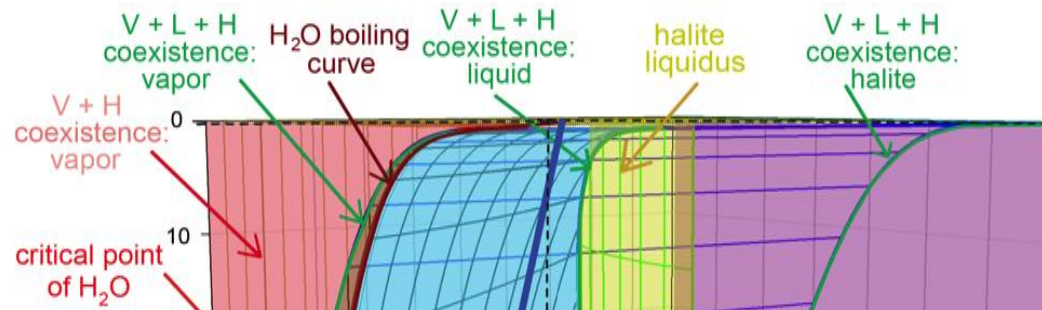
Scott et al., 2017, *Geophysical Research Letters*

Shallow (2-3 km) depth boiling zones with descending hypersaline brines underlain by halite-filled pore space insulate shallow intrusions



Deep condensation zones maximize vapor mass fluxes and overall heat transfer

Intrusion depth controls phase separation style



Thanks for listening!

